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Snowfall brightens Antarctic future

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Abstract

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Snowpacks absorb more sunlight as they warm. The Antarctic Plateau may buck this trend

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over the 21st century since increased snowfall there inhibits the snowpack from dimming.

1 Introduction

The color of snow tells a remarkable story. Of course snow appears white because its reflectance to visible light is uniformly high. However, its reflectance changes with astonishing abruptness at other wavelengths, and is a complex function of the exact ice crystal size and shape^{1,2}. Pristine snow is a valuable shield against global warming that reflects up to 85% of sunlight and traps only the remainder as heat^{1,3,4}. That is why almost imperceptible reductions in snow reflectance due to warming and pollution^{3,5} have become a great concern. Increased heat trapping by darker snow triggers a vicious cycle which amplifies the greying of snow⁵⁻⁷. With temperatures rising globally what, if anything, will oppose the self-reinforced darkening of snow and keep it from melting even faster? On page fxm of this issue, Picard et al.⁸ deduce from snow color measurements that fresh snowfall inhibits the seasonal greying of snow on the Antarctic Plateau by up to 3%, and reduces summertime temperatures there by up to 4 °C. On climate timescales, the increase of Antarctic snowfall expected with 21st century warming may be enough to prevent the surface from further darkening.

Antarctica's reprieve from darker snow would be a welcome surprise because the enemies of snow reflectance are time and temperature, which is projected to rise by about 3 °C this century. Much like ice cubes in a home freezer, snow crystals lose their sharp facets to duller, rounder shapes as they age^{1,9} (Fig. 1). Heat accelerates this metamorphism so that pristine, sharply faceted fresh crystals quickly grow during summer to become larger, rounder aged snow which absorbs more and reflects less sunlight^{1,5,9}. Snow reflectance also changes during wind events (which shatter and sublimate crystals), and due to surface crusts and ripples. Findings reported here suggest these secondary contributors explain less than a third of summer snow reflectance changes on the Antarctic Plateau. Temperature and snowfall are the main players.

The Antarctic Plateau endures long periods of polar night during which its visible reflectance cannot be measured, so Picard and colleagues focused on the seasonal behavior of a reflectance proxy, the snow grain size. First they teased grain size information from the wavelength-dependent surface microwave emissions measured daily by meteorological satellites. A sophisticated model

31 of the microwaves traveling from the surface through the atmosphere best matches the measured
32 signal when the snowpack is modeled as smaller, younger surface grains atop larger, inactive snow
33 grains deposited in previous seasons.

34 To confirm the findings deduced from the satellite imagery required “ground truth” measure-
35 ments of grain sizes in Antarctic snow. For this the authors built an optical probe that operates at
36 infrared wavelengths selected for sensitivity to snow grain size and brought it to Dome C, high atop
37 the Plateau, one of the coldest places on Earth. Daily sampling showed that snow grains nearer the
38 surface grow much faster through the brief summer season than do the deeper snow grains which
39 are insulated from the relatively warm daily maximum surface temperatures of summer.

40 The 10-years of daily satellite measurements, calibrated with model and *in situ* results, show
41 that surface snow on the Antarctic Plateau undergoes a remarkably consistent annual cycle. Snow
42 crystals deposited in the polar night remain small (and therefore bright) because metamorphism
43 is sluggish at winter temperatures that can fall below -80°C . Grains grow once temperatures
44 begin to rise by December, and reach their maximum annual size by February at a balmy -25°C .
45 The stronger growth in years with weaker summer snowfall causes reflectance to drop by a few
46 percent. How exactly does snowfall reduce the reflectance? Not simply by burying large old
47 crystals with small fresh ones, although that helps. Accumulated summer snowfall is usually less
48 than 1 cm thick, but its reflectivity is high enough to chill the underlying crystals and short-circuit
49 their temperature-driven growth (and greying).

50 Snowpack properties such as reflectance are notoriously heterogeneous^{2,10} as comparison to
51 other regions shows. For instance, Greenland’s summertime reflectivity has decreased signifi-
52 cantly over the past decade⁷. Is summertime grain growth and rounding there already proceeding
53 too rapidly to be fully compensated by increased snowfall? Snow accumulation gains partially
54 offset the accelerating loss of Antarctic ice¹¹, and climate models project that snowfall in the inte-
55 rior Plateau will increase with 21st century warming. If that snowfall inhibits surface dimming at
56 present rates, the findings here indicate that the Antarctic Plateau could (unlike Greenland) main-
57 tain its high reflectance.

58 Fresh snow is the brightest surface on Earth, outshining glaciers, sea ice, deserts, and even the
59 thickest clouds, and Picard *et al.* hint at many important questions regarding snow/climate inter-
60 actions. For instance, climate models inadequately represent the snow grain size and shape distri-
61 butions that determine not only the reflectance studied here, but also snow thermal, hydraulic, and
62 mechanical behavior. What exciting effects might these connections have on surface temperature
63 and hydrology? And, given that snow crystal nucleation and metamorphism are poorly understood,
64 how do they alter reflectance in coastal regions of Antarctica and Greenland where snow is much
65 nearer the freezing point and is subject to strong katabatic winds? What about in tundra, alpine,
66 and sub-alpine regions? The authors combined lines of evidence from multi-channel satellite re-
67 mote sensing, *in situ* monitoring, an active field campaign, and snowpack and radiative models.
68 Their findings highlight needed improvements in snow/climate interactions in climate models, and
69 shows Antarctica's future is brighter than previously thought.

70 [Figure 1 about here.]

71 Acknowledgments

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73 from http://dust.ess.uci.edu/ppr/ppr_Zen12.pdf.

74 Bibliography

- 75 1. Warren J. Wiscombe and Stephen G. Warren. A model for the spectral albedo of snow. I: Pure
76 snow. *J. Atmos. Sci.*, 37:2712–2733, 1980.
- 77 2. Stephen R. Hudson and Stephen G. Warren. An explanation for the effect of clouds over snow
78 on the top-of-atmosphere bidirectional reflectance. *J. Geophys. Res.*, 112(D19202), 2007.
- 79 3. Mark Z. Jacobson. The climate response of fossil-fuel and biofuel soot, accounting for soot's
80 feedback to snow and sea ice albedo and emissivity. *J. Geophys. Res.*, 109(D21201), 2004.

- 81 4. Alex Hall and Xin Qu. Using the current seasonal cycle to constrain snow albedo feedback in
82 future climate change. *Geophys. Res. Lett.*, 33(L03502), 2006.
- 83 5. Mark G. Flanner and Charles S. Zender. Linking snowpack microphysics and albedo evolu-
84 tion. *J. Geophys. Res.*, 111(D12208), 2006.
- 85 6. Mark G. Flanner, Charles S. Zender, Peter G. Hess, Natalie M. Mahowald, Thomas H. Painter,
86 V. Ramanathan, and Philip J. Rasch. Springtime warming and reduced snow cover from car-
87 bonaceous particles. *Atmos. Chem. Phys.*, 9(7):2481–2497, 2009.
- 88 7. J. E. Box, X. Fettweis, J. C. Stroeve, M. Tedesco, D. K. Hall, and K. Steffen. Greenland ice
89 sheet albedo feedback: thermodynamics and atmospheric drivers. *The Cryosphere*, 6:821–839,
90 2012.
- 91 8. Ghislain Picard, Florent Domine, Gerhard Krinner, Laurent Arnaud, and Eric Lefebvre. In-
92 hibition of the positive snow-albedo feedback by precipitation in interior Antarctica. *Nature*
93 *Clim. Change*, 6:fxm, 2012.
- 94 9. Florent Dominé, Thomas Lauzier, Axel Cabanes, Loïc Legagneux, Werner F. Kuhs, Kirsten
95 Techmer, and Till Heinrichs. Snow metamorphism as revealed by scanning electron mi-
96 croscopy. *Microsc. Res. Tech.*, 62(1):33–48, 2003.
- 97 10. Xianwei Wang and Charles S. Zender. Arctic and Antarctic diurnal and seasonal variations of
98 snow albedo from multiyear Baseline Surface Radiation Network measurements. *J. Geophys.*
99 *Res. Earth Surf.*, 116(F03008), 2011.
- 100 11. E. Rignot, I. Velicogna, M. R. van den Broeke, A. Monaghan, and J. Lenaerts. Acceleration
101 of the contribution of the Greenland and Antarctic ice sheets to sea level rise. *Geophys. Res.*
102 *Lett.*, 38(L05503), 2011.

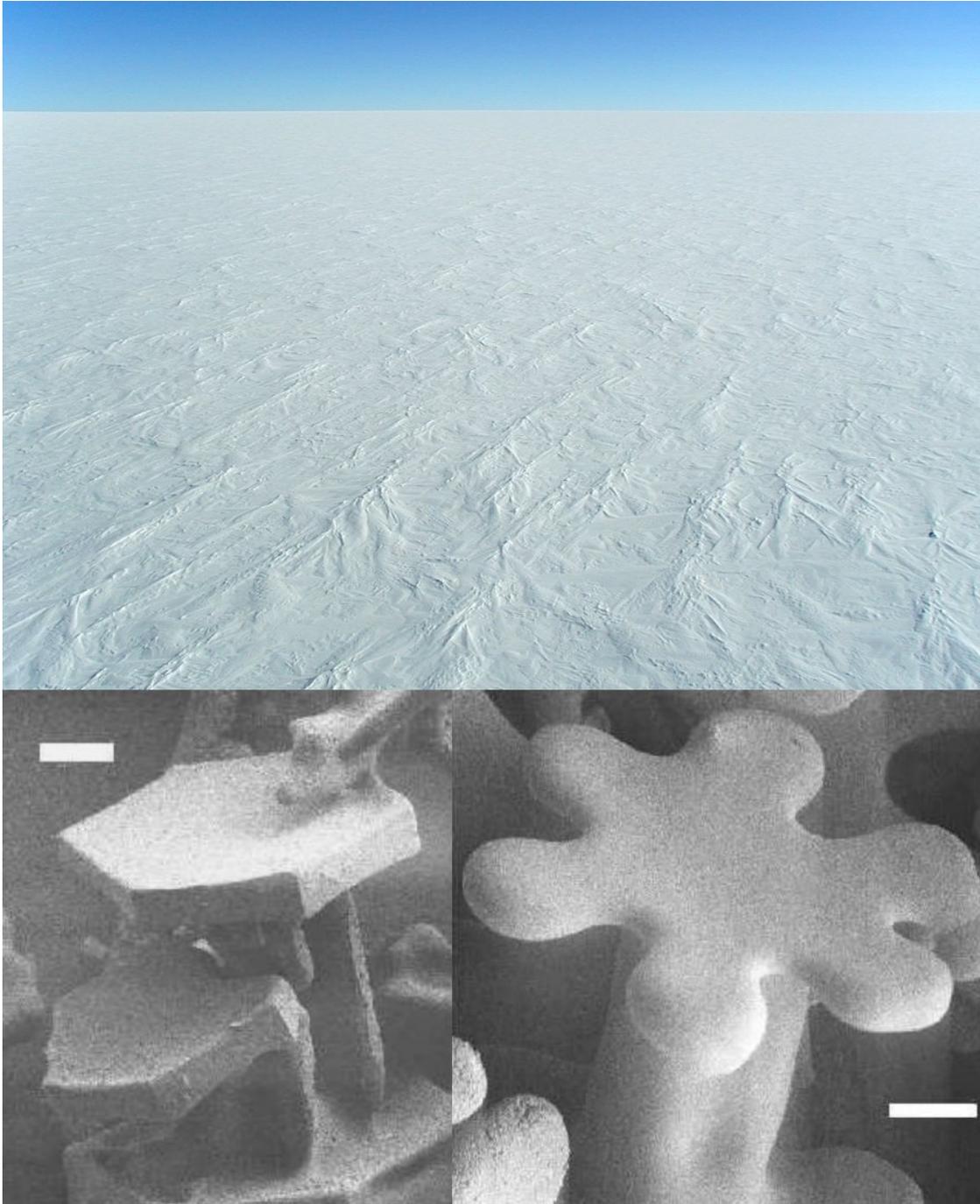


Figure 1: Picard et al.⁸ used near-infrared optics to probe surface snow grain sizes over a summer campaign at **a**, Dome C in Antarctica². The daily *in situ* measurements match changes in grain size retrieved from microwave remote sensing across the high Antarctic Plateau. The warmer temperatures during summer tend to change ice crystals from **b**, the pristine, sharply faceted shapes of fresh snowfall to **c**, the larger, rounder shapes of aged snow which absorbs more and reflects less sunlight^{5,9}. SEM images show metamorphism of Alpine snowpack crystals (scale bars are 100 μm). Daily microwave imagery from 2000–2010 confirm that on seasonal timescales fresh snow offsets much of the grain growth (and thus darkening) due to summer warming. On climate timescales, global warming is expected to increase Antarctic snowfall through the 21st century. This will offset the darkening of the Plateau expected from rounder snow grains in the warmer climate.